## Review of Nd-Sr-Pb isotope for volcanic rock from Ulreung Islands.

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Nd-Sr-Pb isotope data base supports the identification by Hart (1988) of four end member isotope components in oceanic basalts: no EM array (without enriched mantle) or matle plane  $(DMM-HIMU:\Delta Nd=0)$ , hi(gh)-Nd  $(EMII:\Delta Nd=+)$ , lo(w)-Nd(EMI:  $\Delta$ Nd=-). Isotopic ratio of Sr (0.7047-0.7053) and Nd (0.512463-0.512586) for volcanic rock from Ulreung Is. have shown EMI ( $\Delta Nd=-5\sim-7$ ) signature which is similar to Walvis Ridge or Kerguelen. Otherwise, volcanic rocks from Ulreung Is. have shown Dupal anomaly which is similar to Walvis Ridge or Kerguelen. (<sup>206</sup>Pb/  $^{204}$ Pb-  $^{207}$ Pb/  $^{204}$ Pb.  $^{206}$ Pb/  $^{204}$ Pb-  $^{208}$ Pb/  $^{204}$ Pb diagram). Dupal equation is  $\Delta^{206}$ Pb/  $^{204}$ Pb=  $[(^{206}$ Pb/ $^{204}$ Pb)<sub>DS</sub>- $(^{206}$ Pb/ $^{204}$ Pb)<sub>NHRL</sub>] ×100. Here, DS (data set;  $\Delta^{206}$ Pb/ $^{204}$ Pb ratio of the sample for studied area, NHRL(northern hemisphere reference line); that of Mid-Atrantic Ridge, East Pacific Rise, Iceland, Azores, Carnaries Cape Verde. Cheju Is. (Dupal) and SOPITA(south pacific isotopic and thermal anomaly; not Dupal: Samoa) have EMII component. However, SOPITA have HIMU (high<sup>238</sup>U/<sup>204</sup>Pb), Cheju Is. (Dupal) have not HIMU component.  $\Delta^{206}$ Pb/  $^{204}$ Pb ratio for HIMU component is  $-60 \sim -120$ , but  $\Delta^{206}$ Pb/ $^{204}$ Pb ratio for Dupal component is more than +80.  $\Delta^{206}$ Pb/  $^{204}$ Pb ratio for rock from Ulreung Is. is +128~+143. Therefore, rock from Ulreung Is. have not HIMU component. Sr-Pb and Nd-Pb diagram have shown for rock from Ulreung Is. DMM-EMI (mantle) array. Isotope ratio (Nd-Sr-Pb) of Ulreung rock deviated slightly from DMM-EMI. Maybe mantle of magma source for Ulreung Is. is very weakly affected by EMII component.