

RESEARCH ON THE RULE OF SOIL MOISTURE CONTENT DYNAMICS IN ZHALONG WETLAND

WANG XING-JU^{1,4*}, XU SHI-GUO² and WANG HAO³

¹Doctor, School of Civil & Hydraulics Engineering,
Dalian University of Technology, 116-024, PRC
(Tel: +86-411-8470-7680, Fax: +86-411-8470-7680, e-mail: xjw@sdu.edu.cn)

²Professor, School of Civil & Hydraulics Engineering,
Dalian University of Technology, 116-024, PRC
(Tel: +86-411-8470-7680, Fax: +86-411-8470-7680, e-mail: sgxu@dlut.edu.cn)

³Doctor, School of Civil & Hydraulics Engineering,
Dalian University of Technology, 116-024, PRC
(Tel: +86-411-8470-7680, Fax: +86-411-8470-7680,
e-mail: haowang@student.dlut.edu.cn)

⁴School of Civil & Hydraulics Engineering, Shandong University, Jinan, 250-061, PRC

Wetland ecosystems are complex, evolving structures whose characteristics and dynamic properties depend on many interrelated links between climate, soil, and vegetation. On the one hand, climate and soil control vegetation dynamics (e.g. 1, 2, 3, 4, 7, and 20); on the other hand, vegetation exerts important control on the entire water balance and is responsible for many feedbacks to the atmosphere (e.g. 7, 8, 9, 10, 11, 18, and 19). Many important issues depend on the quantitative understanding of this dynamics, including environmental preservation and proper management of resources (e.g. 5, 6, 7, 12, 13, 14, 15, 16, and 17).

Zhalong wetland in north China is a typical reed marsh, and has obviously functions of retaining water and flood control. The soil layers are the base material of the water movement and storage in the wetland. Soil moisture content is the key variable which synthesizes the action of climate, soil, and vegetation on the water balance and the dynamic impact of the water balance on plants. Although other sources of stress (fire, grazing, nutrient availability, etc.) are certainly also present, in the whole wetland ecosystem soil moisture content is the most important resource affecting vegetation structure and organization. Otherwise, in this area, the seasonal frozen soils distribute extensively, which changing the physical appearance of soil moisture content delivers, resulting in the instability of runoff generation and accordant junction making it difficult to simulate the runoff.

The research results show that the soil moisture content, in locations with difference topography, vegetation, and the soil depths, is difference.

Soil moisture content in reed marsh is obviously larger than in farmland and grassland in the same period, and soil moisture content in farmland is slightly larger than in grassland; soil moisture content in reed marsh drastically changes and remarkably undulates with time, while in farmland and grassland, soil moisture content changes gently, especially in farmland; soil moisture content in all locations decreases with the increase of soil depth, besides, decreases at 5-30 cm depth are very sharp, from 30 cm depth, decreases are gently, and stable in grassland.

Soil moisture content heightens with reducing topography, soil moisture content in flat

ground is obviously larger than in sloping field and hillock in the same period, and soil moisture content in sloping field is slightly larger than in hillock; besides, with the increase of soil depth, soil moisture content both in sloping field and hillock tends to be stable, while in flat ground it still changes drastically.

Because of effects due to temperature gradient difference in soils, soil water moves upwards and congregates in freezing interface, as a result, soil moisture content in frozen zones currently increases as much as 20%-40% than in the unfrozen zones in the whole freezing period, as well as, in the low and flat locations, frozen soil moisture content often exceeds field capacity, or even becomes saturation and supersaturation.

REFERENCES

- Wang G.X., *et al.* (2003). Study on the influence of vegetation change on soil moisture cycle in Alpine meadow, *Journal of glaciology and geocryology*, pp. 653-658 (in Chinese).
- Van der Kamp G, *et al.* (1999). Drying out of small prairie wetlands after conversion of their catchments from cultivation to permanent brome grass. *Hydrol Sci J.*, 44(3), pp. 387 – 398.
- Van der Kamp G., Hayashi M. and Gallén D. (2003). Comparing the hydrology of grassed and cultivated atchments in the semi-arid Canadian prairies. *Hydrological Processes*. Vol. 17.
- Christie H W. (1985). Soil and subsoil moisture accumulation due to dryland agriculture in southern Alberta. *Can J Soil Sci.*, pp. 805 – 810.
- Pan X. l., Deng W. *et al.* (2003). Classification of hydrological landscapes of typical wetlands in northeast China and their vulnerability to climate change, *Research of environmental science.* (in Chinese).
- Zhai J. l. He Y. and Deng W. (2002). Wetland functions of the Xianghai natural reserve, *Bulletin of Soil and Water Conservation* (in Chinese).
- Lang H. Q. (1998). Biological diversity of wetland in China. Environment science publishing company in China. pp. 89-93 (in Chinese).
- Bryan D Wheeler (1999). Water and plants in freshwater wetlands, In: Baird A J, *et al* eds *Eco-hydrology*. Landon: Rouledge, pp.128~157.
- Euliss N. H. and Mushet D. M. (1996). Water-level fluctuations in wetlands as a function of landscape condition in the prairie pothole region, *Wetlands*, pp. 587 – 593.
- de Jong E. and Kachanoski R G. (1987). The role of grasslands in hydrology, *Canadian Aquatic Resources*, pp. 213- 215. - 241.
- Price J. S. and Waddington J. M. (2000). Advances in Canadian wetland hydrology and biogeochemistry, *Hydrological Processes*, Vol. 14.
- Poiani, K. A. and Johnson, W. C. (1993). A spatial simulation model of hydrology and vegetation dynamics in semi-permanent prairie wetlands, *Ecological Applications*, pp.279–293.
- Lafleur P M, McCaughey J H, Joiner D. W. *et al.* (1997). Seasonal trends in energy, water, and carbon dioxide fluxes at a Northern boreal wetland, *Journal of Geophysical Research*, 102, pp.29009 - 29020.
- Pook E W, Moore P H R, Hall T, *et al.* (1991). Rainfall interception by trees of pinus radiata and eucalyptus viminalis in a 1300 mm rainfall area of southeastern new South Wales, *Hydrol Process*, pp. 127 – 141.
- Sellers P. J., Randall D. A., Collatz G.J., *et al.* (1996). A revised land surface

- parameterization for atmospheric GCMs Part I Model formation, *Climate*, pp. 676 – 705.
- Zeng N., Shuttleworth J. W., Gash J H. C. (2000). Influence of temporal variability of rainfall on interception loss I Part : Point analysis, *J Hydrol*, , 228(324), pp.228 – 241.
- Aboal J. R., Morales D., Hernandez M., *et al.* (1999). The measurement and modeling of the variation of stemflow in a laurel forest in Tenerife, Canary Islands, *Hydrol*, 221(324). pp. 161 – 175.
- Taniguchi M., Tsujimura T., and Tanaka T. (1996). Significance of stemflow in groundwater recharge I : evaluation of the stemflow contribution to recharge using a mass balance approach, *Hydrol Process.* , 10(1), pp.71 – 80.
- Boyd C. E. (1987). Evapotranspiration/ evaporation (E/ E0) ratios for aquatic plants, *Aquatic Plant Management*, pp.1 – 3.
- Price J. S. (1996). Hydrology and microclimate of a partly restored cutover bog, Quebec, *Hydrol Process.* , pp.1263 – 1272.
- Hou Sh. Y., Chen Q. (2003). Functions of flood control and disaster alleviation in Zhalongwetland, *Journal of Heilongjiang hydro. Academy* (in Chinese).
- GUAN Z. C. (2002). Study on hydrological simulation in cold area [D]. Hohai University (in Chinese).