

## SOME APRIORI ESTIMATES FOR THE QUASI-GEOSTROPHIC EQUATION

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ABSTRACT. We present a new apriori estimates for the surface quasi-geostrophic equation. This apriori estimates give a new blow-up criterion which is different from the known Beale-Kato-Majda type criterion.

### 1. Introduction

We consider the quasi-geostrophic equations in the whole 2-dimensional domain  $\Omega = \mathbb{R}^2$ ,

$$(1.1) \quad (QG) \begin{cases} \frac{\partial \theta}{\partial t} + (v \cdot \nabla)\theta = 0, & v = \nabla^\perp \Lambda^{-1} \theta & \text{in } \Omega \times \mathbb{R}_+ \\ \theta(x, 0) = \theta_0(x) & & \text{in } \Omega \end{cases}$$

where  $\theta$  and  $v$ , respectively, are the surface temperature and the velocity of the flow.  $\Lambda = (-\Delta)^{\frac{1}{2}}$  is the pseudo-differential operator defined in Fourier space by  $\widehat{(-\Delta)^{\frac{1}{2}}u}(\mathbf{k}) = |\mathbf{k}|\hat{u}(\mathbf{k})$  and  $\nabla^\perp$  is the orthogonal derivative operator defined by  $(-\partial_2, \partial_1)$ . The surface quasi-geostrophic equation describes the dynamics of large eddies in the atmosphere and ocean. For the geophysical meaning of the surface quasi-geostrophic equation, see [9]. The main mathematical interest in the surface quasi-geostrophic equation lies in the similarities with the 3D Euler equations.  $\nabla^\perp \theta$  plays the similar role of the vorticity for the 3D Euler equations. This direction of the research was first initiated by Constantin, Majda and Tabak[5]. Weak solutions have been constructed by Resnick[10]. The following Beale-Kato-Majda[1] type blow up criterion for the quasi-geostrophic

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Received October 15, 2007.

2000 Mathematics Subject Classification: 35Q30, 35Q35 .

Key words and phrases: surface quasi-geostrophic equation, blow-up criterion.

equation has been proved by Constantin, Majda and Tabak[5].

$$\limsup_{t \rightarrow T^*} \|\theta(t)\|_{H^m} = \infty \quad \text{if and only if} \quad \int_0^{T^*} \|\nabla^\perp \theta(t)\|_{L^\infty} dt = \infty.$$

This criterion has been refined in [2] using Triebel-Lizorkin space. Hyperbolic saddle collapse blow-up was one of the possible singularity formation scenario for the solutions of the quasi-geostrophic equation. The hyperbolic saddle type scenario for the quasi-geostrophic equation has been excluded by Cordoba[6](see [4] for numerical simulations).

Following the method presented in [3], we have

**THEOREM 1.** *Let  $\theta \in C([0, T]; H^m(\mathbb{R}^2))$  be a classical solution to the 2D quasi-geostrophic equations with  $m > 2$ . Suppose that there exists an absolute constant  $\epsilon_0 > 0$  such that for some  $t_0$  with  $0 \leq t_0 < T$ ,*

$$(1.2) \quad \sup_{t_0 \leq t < T} (T - t) \|\nabla \theta(t)\|_{L^\infty(\mathbb{R}^2)} < \epsilon_0.$$

*Then  $\theta \in C([0, T + \delta]; H^m(\mathbb{R}^2))$  for some  $\delta > 0$ .*

**REMARK 1.** *The Theorem 1 implies that if  $T^*$  is the first time of singularity, then we have the following blow-up rate*

$$\limsup_{t \nearrow T^*} \|\nabla \theta(t)\|_{L^\infty(\mathbb{R}^2)} \geq \frac{\epsilon_0}{T^* - t}.$$

*Our blow-up estimate has an advantage over BKM type criterion[5] in the sense that their estimates cannot exclude the possibility that blow-up rate behaves like  $o((T^* - t)^{-1})$ , e.g.,*

$$\|\nabla \theta(t)\|_{L^\infty(\mathbb{R}^2)} \sim O(1/((T^* - t)|\log(T^* - t)|)),$$

*since  $1/(t \log t)$  is not integrable near origin. In contrast, our estimate (1.2) does not allow such blow-up rate.*

## 2. Proof of Theorem 1

In this section we present the proofs of Theorem 1. The following commutator estimate is useful for the proof of Theorem 1 and the proof of the following proposition can be found in [8](see also [7]). The space  $H^{s,p}$  denotes a subspace of  $L^p(\Omega)$ , equipped with the norm  $\|f\|_{H^{s,p}} = \|\Lambda^s f\|_p$ .

**PROPOSITION 1.** *Suppose that  $s > 0$  and  $p \in (1, \infty)$ . If  $f, g \in \mathcal{S}$ , then*

$$\|\Lambda^s(fg) - f\Lambda^s g\|_p \leq C(\|\nabla f\|_{p_1} \|g\|_{H^{s-1,p_2}} + \|f\|_{H^{s,p_3}} \|g\|_{p_4}),$$

where  $\frac{1}{p_1} + \frac{1}{p_2} = \frac{1}{p_3} + \frac{1}{p_4} = \frac{1}{p}$ .

**Proof of Theorem 1.** We first take  $\nabla^\alpha$  operator on the both sides of quasi-geostrophic equation, where  $\alpha = (\alpha_1, \alpha_2, \alpha_3)$  is a multi-index with  $|\alpha| = \sum_{i=1}^3 \alpha_i \leq m$ . Let  $t < T$ . Multiplying (1.1) by  $(T-t)\nabla^\alpha \theta(t)$ , integrating over  $\mathbb{R}^2$ , and summing over  $\alpha$  for  $|\alpha| \leq m$ , we have

$$\frac{1}{2} \frac{d}{dt} \left( (T-t) \|\theta\|_{H^m}^2 \right) + \frac{1}{2} \|\theta\|_{H^m}^2 = -(T-t) \sum_{|\alpha| \leq m} \int_{\mathbb{R}^2} \nabla^\alpha ((v \cdot \nabla) \theta) \nabla^\alpha \theta dx := RHS.$$

Due to the commutator estimates Proposition 1, the righthand side is estimated as follows:

$$(2.3) \quad RHS \leq C(T-t) (\|\Lambda v(t)\|_\infty \|\theta(t)\|_{H^m}^2 + \|\nabla \theta(t)\|_\infty \|v(t)\|_{H^m} \|\theta(t)\|_{H^m}).$$

Since we have  $\Lambda v(t) = \nabla^\perp \theta(t)$  and  $\|v(t)\|_{H^m} = \|\theta(t)\|_{H^m}$ , (2.3) reduces to the following

$$(2.4) \quad RHS \leq C(T-t) \|\nabla \theta(t)\|_\infty \|\theta(t)\|_{H^m}^2.$$

Thus we have

$$\frac{1}{2} \frac{d}{dt} \left( (T-t) \|\theta\|_{H^m}^2 \right) + \left( \frac{1}{2} - C(T-t) \|\nabla \theta(t)\|_\infty \right) \|\theta(t)\|_{H^m}^2 \leq 0.$$

We choose  $\epsilon_0 = \frac{1}{4C}$ , where  $C$  is the absolute constant in (2.4). It is straightforward, from the assumption (1.2), that

$$\frac{1}{2} \frac{d}{dt} \left( (T-t) \|\theta\|_{H^m}^2 \right) + \frac{1}{4} \|\theta\|_{H^m}^2 \leq 0.$$

Therefore, integrating in time from  $t_0$  to  $\tau$  for any  $\tau \in (t_0, T)$ , we obtain

$$(2.5) \quad \sup_{t_0 \leq t < T} (T-t) \|\theta\|_{H^m}^2 + \frac{1}{2} \int_{t_0}^T \|\theta\|_{H^m}^2 dt \leq (T-t_0) \|\theta(t_0)\|_{H^m}^2.$$

Since  $\int_{t_0}^T \|\theta\|_{H^m}^2 dt$  is finite, the conclusion is immediate from BKM type criterion.  $\square$

## References

- [1] J.T. Beale, T. Kato and A. Majda, *Remarks on the breakdown of smooth solutions for the 3-D Euler equations*, Comm. Math. Phys., **94**, (1984), pp. 61–66.
- [2] D. Chae, *The quasi-geostrophic equation in Triebel-Lizorkin space*, Nonlinearity, **16** (2003) pp. 479–495.
- [3] D. Chae, K. Kang, and J. Lee, *Notes on the asymptotically self-similar singularities in the Euler and the Navier-Stokes equations*, preprint (2007).
- [4] P. Constantin, Q. Nie, and N. Schorghofer, *Nonsingular surface quasi-geostrophic flows*, Physics Letters A **241**, (1998), pp. 168–172.
- [5] P. Constantin, A. J. Majda and E. G. Tabak, *Formation of strong fronts in the 2D quasi-geostrophic thermal active scale*, Nonlinearity **7**, (1994), pp. 1495–1533.
- [6] D. Cordoba, *Nonexistence of simple hyperbolic blow-up for the quasi-geostrophic equation*, Ann. of Math. **148** pp. 1135–1152.
- [7] N. Ju, *Existence and uniqueness of the solution to the dissipative 2D quasi-geostrophic equations in the Sobolev space*, Comm. Math. Phys. **251** (2004), pp. 365–376.
- [8] C. Kenig, G. Ponce, and L. Vega, *Well-posedness of the initial value problem for the Korteweg-De Vries equation*, J. Am. Math. Soc., **4** (1991), pp. 323–347.
- [9] J. Pedlosky, *Geophysical fluid dynamics*, New-York: Springer-Verlag, 1987.
- [10] S. Resnick, *Dynamical problems in nonlinear advective partial differential equations*, PH.D. Thesis, University of Chicago, 1995

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