Article

Future subsurface drainage in the light of climate change in Daegu, South Korea

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Abstract

Over the last century, drainage systems have become an integral component of agriculture. Climate observations and experiments using general circulation models suggest an intensification of the hydrologic cycle due to climate change. This study presents hydrologic simulations assessing the potential impact of climate change on subsurface drainage in Daegu, Republic of Korea. Historical and Long Ashton Research Station weather generator perturbed future climate data from 15 general circulation models for a field in Daegu were run into a water management simulation model, DRAINMOD. The trends and variability in rainfall and Soil Excess Water (SEW) were assessed from 1960 to 2100. Rainfall amount and intensity were predicted to increase in the future. The predicted annual subsurface drainage flow varied from -35 to 40 % of the baseline value while the SEW varied from -50 to 100%. The expected increases in subsurface drainage outflow require that more attention be given to soil and water conservation practices.

Keywords subsurface drainage, climate change, general circulation model, DRAINMOD

Introduction

A well-designed subsurface drainage system with reasonable drain space and depth contributes to large ratio of desalination and high crop yield (Shao et al. 2012). In the Republic of Korea, subsurface drainage has been implemented in 13% of the wetlands to control water logging and land salinization (Jung et al. 2010). The subsurface drainage project sites include Buyeo, Dongjin and Hanan districts among others (Kim and Goo 1977). The hydrology of fields with single or non-parallel drains may be simulated by determining effective drain spacing by calibration (Skaggs et al. 2012). Most of the existing drainage systems were designed according to the American Society of Agricultural and Biological Engineers (ASABE) scientific criteria and with drain spacing of 7 to 15 m and depths of 0.5 m (Jung et al. 2010). Subsurface drainage is a function of local site conditions including climate, soil, cropping system, farming practices, and drainage system (Skaggs et al. 2012a). Approaches to subsurface drainage engineering have assumed the stationarity of rainfall series. Comprehending the response of precipitation to climate change assists in climate change mitigation and adaptation interventions on subsurface drainage (Coulibaly and Shi 2005).

Analyses of historical data trends have shown evidence of temporal changes in hydro-climatic variables (Jung et al. 2011). Examination of observed daily precipitation data from the Republic of Korea showed increasing trends in the summer precipitation amount and intensity (Chang and Kwon 2007). It is widely recognized that in Korea, the impacts of climate change on subsurface drainage will manifest more through changes in extremes than as a result of changes in the mean climate (Xu et al. 2012). Generally, the overall consensus amongst GCM predictions and observations of historical climate data is consistent with monotonic change for temperature (Rogelj et al. 2012).

Difficulties in modelled extreme rainfall result from a lack of enough data to provide a stable estimate of their frequency and intensity. Nevertheless, many studies have shown that the projections from GCMs are indicative of what we may expect from future rainfall extremes (Fowler et al. 2010). Review of pertinent literature shows that the impacts of climate change...
on sub surface drainage in Daegu have not been previously studied. The objective of this study is to simulate the impact of climate change on the sub surface drainage systems currently installed in vicinity of Daegu.

Study Area

Daegu lies in a basin surrounded by low mountains. The Geumho River flows along Daegu’s northern eastern boundary, emptying in the Nakdong River. The soils in Daegu were identified to be gray shale which consists of 2:1 minerals like illite and vermiculite and were derived from parent material residuum (Um et al. 1993). Land use in Daegu can be classified into water bodies (2%), irrigated crops (17%), forests (57%), grasslands (12%) and urban area (12%) (Lee and Kim 2008).

Daegu’s climate is humid subtropical climate with an average annual temperature of 13.7°C, the average temperature in August is the hottest 26.1°C and the coldest 0.2°C in January. Average annual rainfall is only 1027.9 mm.

Methods

While we do not attempt to perform a rigorous detection and attribution study, changes in the key parameters with time were investigated using simple statistical methods.

Climate Change Data

Historical data from 1960 to 1990 was extracted from the Korean Meteorological Administration (KMA) (www.kma.go.kr) and was adopted as the baseline in this study. Future climate change scenarios for 2011-2030 (2020s), 2045-2065 (2055s) and 2080-2100 (2090s) and A2, A1B and B1 Special report on emissions scenarios (SRES) scenarios were generated stochastically by perturbing the baseline climate in line with the outputs from a 15 GCMs using LARS-WG (Long Ashton Research Station stochastic Weather Generator). LARS-WG is computationally inexpensive and enables the efficient production of large ensembles of scenarios. The LARS-WG model simulates precipitation occurrence using two - state, first order Markov chains: precipitation amounts on wet days using the gamma distribution; temperature and radiation components

Table 1 GCMs information and predicted changes in temperature and rainfall by the 2090s

<table>
<thead>
<tr>
<th>Centre</th>
<th>GCM Acronym</th>
<th>Predicted ΔT (°C)</th>
<th>ΔR (%)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Bjerknes Centre for Climate Research, Norway</td>
<td>BCM2</td>
<td>3.0(19.8)</td>
<td>1.8(8.9)</td>
</tr>
<tr>
<td>Canadian Centre for Climate Modelling and Analysis,</td>
<td>CGMR</td>
<td>2.9(16.9)</td>
<td></td>
</tr>
<tr>
<td>Centre National de Recherches Meteorologiques, France</td>
<td>CNCM3</td>
<td>4.2(16.7)</td>
<td></td>
</tr>
<tr>
<td>Australia’s Commonwealth Scientific and Industrial Research Organization</td>
<td>CSKMK3</td>
<td>2.4(18.6)</td>
<td>1.6(4.0)</td>
</tr>
<tr>
<td>Institute of Atmospheric Physics, China</td>
<td>FGOALS</td>
<td>2.1(15.5)</td>
<td>1.3(54)</td>
</tr>
<tr>
<td>Geophysical Fluid Dynamics Laboratory, USA</td>
<td>GFCM21</td>
<td>3.7(2.1)</td>
<td>3.5(3.1)</td>
</tr>
<tr>
<td>Goddard Institute for Space Studies, USA</td>
<td>GIAOM</td>
<td>2.3(2.8)</td>
<td>1.9(1.5)</td>
</tr>
<tr>
<td>UK Met. Office</td>
<td>HADCM3</td>
<td>4.4(45.2)</td>
<td>4.4(21.7)</td>
</tr>
<tr>
<td>Institute for Numerical Mathematics, Russia</td>
<td>INMCM3</td>
<td>4.0(17.5)</td>
<td>3.2(10.6)</td>
</tr>
<tr>
<td>Institute Pierre Simon Laplace, France</td>
<td>IPCM4</td>
<td>4.2(15.9)</td>
<td>4.1(14.8)</td>
</tr>
<tr>
<td>National Institute for Environmental Studies, Japan</td>
<td>MIHR</td>
<td>4.5(24.1)</td>
<td>3.4(21.6)</td>
</tr>
<tr>
<td>Max Planck Institute for Meteorology, Germany</td>
<td>MPEH5</td>
<td>3.9(5.2)</td>
<td>3.5(10.2)</td>
</tr>
<tr>
<td>National Centre for Atmospheric Research, USA</td>
<td>NCCCSM</td>
<td>4.0(17.8)</td>
<td></td>
</tr>
</tbody>
</table>

using first-order trivariate autoregression that is conditional on precipitation occurrence (Semenov et al. 1997). Table 1 shows the GCMs used in this study and summarizes the projected annual changes from the baseline in rainfall and ambient temperature across the 15 member GCM ensemble by the 2090s.

Correlation of rainfall to subsurface drainage

The subsurface drainage response of a given soil system is governed by soil type, agricultural management practices, rainfall patterns, topography and subsurface conditions (Singh et al. 1996). The correlations of rainfall to subsurface drainage are investigated here because they can indicate a predictive relationship that can be exploited in simulating the response of drainage systems to climate change.

Detection of Rainfall Trends

Analysis of variance (ANOVA) is one of the statistical methods used to determine the differences in different data sets. The single factor ANOVA test was used to determine if there are at least two population means significantly different within each time slice and SRES scenario.

Maximum Rainfall

Annual maxima of daily rainfall for the years 1961-2001 were modelled for five locations in South Korea and there was no evidence suggesting trends in the raw data (Nadaraja and Choi 2007). It is against this background that rainfall return periods were used in order to assess the extreme rainfall. The different rainfall series were fitted to numerous statistical distributions and the most suitable was selected from its ranking by the Kolmogorov-Smirnov and Anderson Darling test. The 3 parameter Gamma distribution (equation 1) was then selected as a statistical model which will capture the difference in behaviour between the baseline and the LARS-WG simulated future scenarios. The shape parameter (\( \alpha \)), scale parameter (\( \beta \)) and location parameter (\( \gamma \)) were estimated from the annual daily rainfall maxima series for the respective time slice for the baseline and for each GCM for the future scenarios.

\[
f(x) = \frac{1}{\beta^\alpha \Gamma(\alpha)} (x - \gamma)^{\alpha - 1} e^{-(x - \gamma)/\beta}; \ x \geq \gamma
\]  

The appropriate 2, 5, 10, 20, 50 and 100 year return period rainfall was calculated from the 3 parameter Gamma survival function.

Mean Rainfall

It has already established that there are no apparent trends in the raw annual mean rainfall (Section 3.2.1) because of the high natural climate variability, therefore the analysis of the trends in mean rainfall were investigated over longer periods of time to deduce the impact of climate change.

Drainage modelling

DRAINMOD is a water management simulation model, which was developed for analysis of soil water movement on a field scale (Skaggs 1980). DRAINMOD is one of the most applied models for the design and evaluation of water management systems (Borin et al. 2000) and was selected for this study. The model is based on the assumption that lateral water movement occurs mainly in the saturated region, drainage flow is computed by using either the Hooghoudt equation (Eq. 2) or the Kikham equation (Eq. 3). This approach assumes an elliptical water table shape and is based on the Dupuit-Forchheimer assumptions with corrections for convergence near the drain lines.

\[
Q = \frac{8kd, m + 4kL}{L^2} \quad (2)
\]

\[
Q = \frac{4\pi k(t + b - r)}{GL} \quad (3)
\]

where Q is the drainage discharge, \( d \), is the equivalent depth, \( m \) is the midpoint water table height above the drain, \( k \) is the lateral saturated hydraulic conductivity, \( L \) is the distance between drains, \( t \) is the surface water depth, \( b \) is the depth from drain to the surface, \( r \) is drain tube radius. \( G \) is a function of \( L \), \( r \), \( d \) (the depth from the drain to the impermeable layer) and \( h \) (the depth from the surface to the impermeable layer) as given in equation 4 below:

\[
G = 2h \left[ \frac{\sin(\pi d/r) - (\pi d/r)}{\tan(\pi d/r)} \right] + \left[ \frac{\pi h}{2} \right] \left[ \frac{\cos(\pi d/r) + \cos(\pi b/r)}{\cos(\pi d/r) - \cos(\pi b/r)} \right] \left[ \frac{\cos(\pi d/r) - \cos(\pi b/r)}{\cos(\pi d/r) + \cos(\pi b/r)} \right] \]  

(4)

The primary input data required to run DRAINMOD include weather data, soil data, crop data and drainage system parameters. The parameters in the above equations are shown in Fig 2.
For soil input data, DRAINMOD requires the relationship between the water table depth and each of upward flux and drainage volume, the soil water characteristic curve, Green-Ampt parameters, lateral saturated hydraulic conductivity. Input data for the model were obtained from the Korean Meteorological Administration (www.kma.go.kr), field experiments and estimated from values presented in literature on nearby or within the study areas (Kang et al. 2002, Nkomoepi and Chung, 2011, http://clic.csos.vt.edu/comanth/06-Asia_Data.pdf) due to time and cost limitations. Results from Borin et al. (2006) indicated that even very limited input data (texture and porosity of the top 30 cm of soil) gave good predictions. Typical gridiron subsurface drainage systems were simulated. A summary of the input drainage design parameters is given in Table 2. In this study the subsurface drainage performance criteria was the Soil Excess Water (SEWd) which is a measure of stress caused by excessive soil water in the top 30 cm.

### Table 2 Summary of input drainage design parameters

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td>Drain depth (cm)</td>
<td>50 – 110</td>
</tr>
<tr>
<td>Drain spacing (cm)</td>
<td>700 – 1,500</td>
</tr>
<tr>
<td>Effective radius of drains (cm)</td>
<td>3</td>
</tr>
<tr>
<td>Drainage coefficient (cm/d)</td>
<td>5</td>
</tr>
<tr>
<td>Actual distance from surface to impermeable layer (cm)</td>
<td>215</td>
</tr>
<tr>
<td>Initial water table depth (cm)</td>
<td>50</td>
</tr>
</tbody>
</table>

Results and Discussion

**Correlation of rainfall to drainage**

Climatic variables such as rainfall in particular determine the hydrology of subsurface drained landscapes (Singh et al. 2009). Fig 3 shows the correlation between baseline rainfall and the simulated subsurface drainage. A strong correlation ($R^2 = 0.87$) exists between rainfall and subsurface drainage. It is evident that subsurface drainage will be sensitive to changes in the amount and intensity of rainfall. The environment is also vulnerable to changes in drainage out flow because subsurface drainage has been identified as a major salt exporter from irrigated areas (Wabba and Christen 2006).

**Fig 3. Correlation between baseline rainfall and the simulated subsurface drainage**

**Table 3 ANOVA tests on rainfall data**

<table>
<thead>
<tr>
<th>Time Slice</th>
<th>SRES Emissions Scenario</th>
<th>Degree of freedom</th>
<th>F value</th>
<th>p value</th>
</tr>
</thead>
<tbody>
<tr>
<td>2020s</td>
<td>A2</td>
<td>8</td>
<td>7.70</td>
<td>&lt;0.0001</td>
</tr>
<tr>
<td></td>
<td>A1B</td>
<td>12</td>
<td>2.51</td>
<td>0.003</td>
</tr>
<tr>
<td></td>
<td>B1</td>
<td>9</td>
<td>1.39</td>
<td>0.186</td>
</tr>
<tr>
<td>2050s</td>
<td>A2</td>
<td>8</td>
<td>4.88</td>
<td>&lt;0.0001</td>
</tr>
<tr>
<td></td>
<td>A1B</td>
<td>12</td>
<td>3.50</td>
<td>&lt;0.0001</td>
</tr>
<tr>
<td></td>
<td>B1</td>
<td>9</td>
<td>2.24</td>
<td>0.017</td>
</tr>
<tr>
<td>2090s</td>
<td>A2</td>
<td>6</td>
<td>8.55</td>
<td>&lt;0.0001</td>
</tr>
<tr>
<td></td>
<td>A1B</td>
<td>10</td>
<td>2.79</td>
<td>0.002</td>
</tr>
<tr>
<td></td>
<td>B1</td>
<td>9</td>
<td>4.39</td>
<td>&lt;0.0001</td>
</tr>
</tbody>
</table>

Trends in rainfall

This study used large rainfall data sets from 15 GCMs. The single ANOVA tests results from each time slice and scenario are shown in Table 3. At least 2 means from the majority of the different rainfall series are established to be significantly ($p < 0.01$) different at the 0.05 alpha level. The 2020s B1 and 2050s rainfall series were not significantly different. This might be due to the fact that the B1 is the least emissions scenario therefore with little changes to the data they remain close to the baseline values from which all the data is perturbed from.
Mean Rainfall
There is general agreement in the predicted mean long term data of the 15 GCMs. Mean rainfall is predicted to drop in the 2020s but gradually increase in the 2055s and 2090s. Table 1 shows some of the predicted changes in rainfall by the 2090s. A majority of the models predict significant increases in rainfall with the exception of B1, the lower emissions scenario. The harmony between analyses of baseline data and multiple GCM data used herein suggests that the increase in rainfall is consistent with the impacts of climate change rather than mere natural climate variability.

Maximum daily Rainfall
The 15 GCMs unanimously predicted increasing amounts of maximum daily rainfall corresponding to the given return period as shown in Fig 4. The 50 and 100 year return period maximum daily rainfall was predicted to be as large as double for the A2 scenario as shown in Fig 4. The increase in rainfall intensity could compromise the functionality of the current drainage systems therefore increasing failure risks (Pu et al. in press). There variability (interquartile range) increases as the return period in increases. The range also generally increases in the future time slices. In this section, we dealt with the maximum daily precipitation amount because of limitations in data availability. It would be more relevant to use 4, 6 or 12 hour maximum rainfall amounts which are more applicable in DRAINMOD.

DRAINMOD SIMULATIONS

Baseline
SEW₀ values are known to be dependent on both surface and subsurface drainage (Skaggs 1980a). Fig 5 shows simulated SEW₀ values for the different combinations of drain spacing and depth. The marginal benefit of drain spacing decreases with increase in the drain depth. It is apparent that in any case wider spacing or deeper depth will decrease the SEW₀. Other fundamental principles including those of ecology, economics and agronomy also have to be considered in the selection of the suitable drain depth, spacing etc. (Cuenca 1989). For example, selecting shallower drain depth does not deliver major reduction in implementation costs but savings would be more environmental. Salt from deep subsoil is not disturbed, resulting in lower salt loads in the drainage effluent (Van Schilfgaarde 1974). Combinations of 4 drainage systems comprising of 50 and 90 cm drain depth and drain spacing of 10 and 15m were selected to be modelled for the future time slices.

Detailed discussions of drain design specifics are sacrificed for the sake of other considerations in this paper. Most relevant to this paper is how the future subsurface drainage will compare to the baseline in the light of climate change.

![Fig 4. Predicted future maximum daily rainfall intensity with respect to the baseline for different return periods](image-url)
Future subsurface drainage in Daegu

Fig 5. Simulated SEW$_{30}$ for various drain spacing and depth in the baseline period.
(50, 70, 90 and 110 refer to drain depth in cm).

Future scenarios
The spread of the simulated subsurface drainage and SEW$_{30}$ of the 15 GCMs are given in box and whisker plots shown in Fig 6. The box represents the interquartile range and the error bars represent the minima and maxima. The line in the boxes represents the mean. There is not sufficient evidence available to conclude a consistent trend in both SEW$_{30}$ and drainage discharge. The simulations do not allude to any consensus as is evident by the means, large ranges and interquartile ranges in Fig 6. For the SRES A2 and A1B (with the exception of 2090s A1B) scenarios, SEW$_{30}$ values exhibit similar skewness patterns, where in the baseline the distributions are close to symmetric in the 2020s but tend to gravitate towards left skewness in the future. The B1 however shifts from left

Fig 6. Simulated future SEW$_{30}$ and subsurface drainage variation with respect to those in baseline period.
skewness in the 2020s towards a symmetric distribution in the future. Generally, the maximum SEW_{t0} will decrease from the baseline values with the highest values emanating from the A2 scenario indicating lower drainability in the future. In the most extreme case, the SEW_{t0} was simulated to have almost doubled by the 2090s. The predicted annual subsurface drainage flow varied from 35 to 40 % of the baseline value while the SEW_{t0} varied from -50 to 100%. Hydrological consequences of this extent of climate change impact could have a major impact on the design of drainage systems and other hydraulic works in Daegu Drainage discharge and SEW_{t0} varied mostly for the B1 and least in the A2 scenario. Drainage values will increase from the baseline values in the 2020s and 2055s but decrease in the 2090s. Despite the changes in mean and maximum daily rainfall being distinctly perceptible, the changes in drainage and SEW possess ambiguity. This could be attributed to the predicted increases in temperature and therefore ET.

Of the selected drainage systems, the 15m spacing and 0.5m depth (15, 0.5) system showed the highest mean relative changes in SEW_{t0} and were the most vulnerable to climate change. The 10m spacing and 0.9m depth (10, 0.9) system showed the least relative changes to climate change. The simulation results revealed a weakness of the multi-GCM approach in which the system with the best mean performance has worst extreme. For example the 10, 0.9 system has the least mean relative changes in SEW_{t0} but also has the highest extreme change (maximum). It should be noted that the average relative changes should not be interpreted as some measure of consensus or a robust statistic from the 15 GCMs but as a statistical measure of central tendency. While the selection of a single GCM is a major source of uncertainty, the results of multiple GCM simulations are not easily conveyed to end users.

Conclusion
In this paper, the impact of climate change on subsurface drainage system and on the amount and intensity of rainfall in Daegu area has been assessed. It is shown that long term rainfall will decrease in the 2020s and increase in the 2055s and 2090s. The likelihood of occurrence of extreme rainfall events will increase particularly for the 50 and 100 year return period maximum daily rainfall. The drain outflow and SEW_{t0} were also predicted to increase in the future. The SEW_{t0} was found to increase by almost 100% in some cases. Hydrological consequences of this extent of climate change impact could have a major impact on the design of drainage systems and other hydraulic works in Daegu area. The assessment of the potential impact of climate change on water systems should be a part of hydrological research in order to implement the appropriate climate change mitigation and adoption measures.

References


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